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Auroral Electron acceleration by Alfvén waves and electrostatic fields

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Abstract. We present a two-dimensional numerical model for the formation of discrete auroral arcs. This model describes the evolution of shear Alfvén waves generated by a growing force near the equatorial plane, and the transition to electrostatic fields when the force becomes stationary. The parallel electric fields on auroral field lines may be regarded as shear Alfvén waves driven by a magnetospheric generator at zero frequency. In our collisionless model, precipitating auroral electrons are accelerated to an energy of 350 eV when the upward current is $3.1 \mu\text{Am}^{-2}$. We also find that the electrostatic potential drop is proportional to the square of the current density.

1. Introduction

Discrete auroral arcs are perhaps the most spectacular phenomenon that regularly can be observed on the night sky. They have been studied intensely by space scientists for several decades, but as indicated in a review of 22 different theories for auroral arcs by *Borovsky* [1993], our understanding of the mechanisms behind the formation of discrete auroral arcs is still rather fragmented and uncertain.

Discrete arcs, in contrast to diffuse aurora, are characterized by the precipitation of field-aligned electrons with a rather well defined peak in their energy spectrum, and it is generally agreed that these electrons have been accelerated by a parallel electric field. Since the magnetospheric plasma is collisionless its parallel conductivity is expected to be very high, and any parallel electric fields should be short-circuited. Several theories have been proposed to explain how parallel potential drops reaching up to 10 kV can exist along auroral field lines, in spite of the high conductivity. Anomalous resistivity caused by wave turbulence [e.g., *Papadopoulos*, 1977] and electrostatic double layers [e.g., *Block*, 1972] have been suggested, but much of the discussion has recently revolved around two main groups of models for the parallel electric fields. One of these groups is purely electrostatic models, such as the kinetic models in which the magnetic mirror force prevents short-circuiting of the potential drop [*Knight*, 1973; *Fridman and Lemaire*, 1980; *Janhunen and Olserson*, 1998], and self-consistent fluid models [*Mitchell et al.*, 1992; *Ganguli et al.*, 1993; *Rönmark*, 1999]. In the other group the parallel electric field associated with

shear Alfvén waves is invoked to accelerate the auroral electrons [*Goertz and Boswell*, 1979; *Haerendel*, 1983; *Lysak and Dum*, 1983; *Lysak*, 1991; *Kletzing*, 1994; *Streltsov and Lotko*, 1995; 1999; *Génot et al.*, 1999; *Streltsov*, 1999]. The model presented here combines features from these two groups into a common framework.

In this study we argue that the auroral display is powered by mechanical forces in the outer magnetosphere. These forces result from the braking of earthward plasma flows and the pressure gradients and anisotropies built up in this process. In a stationary state the forces drive an MHD generator that generates field-aligned currents to and from the auroral ionosphere [e.g., *Birn et al.*, 1999]. Information about changes in the driving force is transmitted by shear Alfvén waves that adjust the field-aligned currents. To carry the upgoing current from the ionosphere, downgoing electrons are accelerated to keV energies [*Rönmark*, 1999], and the precipitation of these accelerated electrons causes the discrete aurora.

In our model a prescribed force is applied to the equatorial part of an auroral flux tube that initially is in static equilibrium. The equations defining the model are first derived, and the boundary conditions and parameter values used in this study are specified. In section 3 we follow the dynamic response of the flux tube in terms of electromagnetic fields, currents, and electron acceleration. Finally, we discuss some properties of the model and its implications for our understanding of auroral arc formation. Some numerical details are treated in Appendix A.

2. The Auroral Circuit Model

2.1. Model Equations

The geometry of the auroral current circuit and the generator region in the equatorial magnetosphere is

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Figure 1. The geometry of the auroral current circuit and the generator region in the equatorial magnetosphere.

sketched in Figure 1. To derive our numerical model we introduce a coordinate system based on the flux tube geometry, with z along the magnetic field lines, x pointing to the earth and y increasing to the west. The origin of these coordinates is at a point where an auroral field line crosses the equatorial plane. Since x and y are constant along field lines, a distance dl is related to these coordinates by the metric

$$dl^2 = (dx^2 + dy^2) \frac{B_0}{B_z} + dz^2, \quad (1)$$

where $B_z(z)$ is the local, time-independent, magnetic field and B_0 is the field strength at $z = 0$ in the equatorial plane. We neglect the curvature of the field lines, and assume homogeneity in the y direction. The current density $\mathbf{j} = j_x \hat{\mathbf{x}} + j_z \hat{\mathbf{z}}$ will then according to Maxwell's equations couple to the electric field components E_x , E_z , and the magnetic field B_y . Ampere's and Faraday's laws for these fields are in our coordinate system expressed as

$$\begin{aligned} \partial_t E_x &= -c^2 \left(\partial_z B_y - B_y \frac{\partial_z B_z}{2B_z} + \mu_0 j_x \right), \\ \partial_t E_z &= c^2 \left(\sqrt{\frac{B_z}{B_0}} \partial_x B_y - \mu_0 j_z \right), \\ \partial_t B_y &= \sqrt{\frac{B_z}{B_0}} \partial_x E_z - \partial_z E_x + E_x \frac{\partial_z B_z}{2B_z}. \end{aligned} \quad (2)$$

We use two fluids to describe the electrons, and let the subscript $s = M$ denote magnetospheric electrons while $s = I$ denotes the ionospheric population. The perpendicular motion of the electrons is given by the $\mathbf{E} \times \mathbf{B}$ drift and can be neglected, since it gives no contribution to the current. The motion along the field lines of an electron component with density n_s and velocity v_s is described by the continuity equation

$$\partial_t n_s = -\partial_z n_s v_s + n_s v_s \frac{\partial_z B_z}{B_z} \quad (3)$$

and the equation of motion. We assume that in a static equilibrium with density n_{0s} and temperature T_s , the

pressure gradients $\partial_z n_{0s} T_s$ are balanced by external forces F_{0s} , that is, $F_{0s} = \partial_z n_{0s} T_s$. The electron velocity v_s is then determined by the equation

$$\partial_t v_s = -v_s \partial_z v_s - \frac{e}{m_e} E_z - \frac{\partial_z n_s T_s}{n_s m_e} + \frac{F_{0s}}{n_{0s} m_e}. \quad (4)$$

The ion motion along the field lines is neglected, and we assume that the ion velocity u_y in the y direction is given by the $\mathbf{E} \times \mathbf{B}$ drift. We will consider the ions to be cold, which implies that the kinetic effects on Alfvén wave propagation are neglected, although they may be significant at altitudes above a few R_E . Including a prescribed mechanical force F_y that drives the generator, the y component of the equation of motion for the ions is

$$d_t u_y = -\frac{e}{m_i} u_x B_z + \frac{F_y}{n_i m_i}. \quad (5)$$

Inserting $u_y = -E_x/B_z$ allows us to calculate the ion velocity u_x in the x direction as

$$u_x = \frac{m_i}{e B_z^2} \frac{\partial_t E_x + \frac{B_z}{n_i m_i} F_y}{1 - \frac{m_i}{e B_z^2} \sqrt{\frac{B_z}{B_0}} \partial_x E_x}. \quad (6)$$

The perpendicular current, which essentially is the ion polarization current, can then be written as

$$\mu_0 j_x = V_A^{-2} \frac{\partial_t E_x + \frac{B_z}{n_i m_i} F_y}{1 - \frac{m_i}{e B_z^2} \sqrt{\frac{B_z}{B_0}} \partial_x E_x}, \quad (7)$$

where $V_A = B_z / \sqrt{\mu_0 n_i m_i}$ is the Alfvén velocity. Inserting this in Ampere's law, we find

$$\partial_t E_x = -A^2 c^2 \left(\partial_z B_y - B_y \frac{\partial_z B_z}{2B_z} \right) - (1 - A^2) \frac{B_z}{n_i m_i} F_y, \quad (8)$$

where

$$A^2 = V_A^2 \frac{1 - \frac{m_i}{e B_z^2} \sqrt{\frac{B_z}{B_0}} \partial_x E_x}{c^2 + V_A^2 (1 - \frac{m_i}{e B_z^2} \sqrt{\frac{B_z}{B_0}} \partial_x E_x)}. \quad (9)$$

We now introduce new dimensionless independent variables through

$$x = \frac{\Omega_i}{c} x, \quad z = \frac{\Omega_i}{c} z, \quad t = \Omega_i t, \quad (10)$$

where $\Omega_i = e B_0 / m_i$, and the dimensionless fields

$$\begin{aligned} E_x &= E_x / \sqrt{c^2 B_0 B_z}, \\ E_z &= E_z / (c B_0), \\ B &= B_y / \sqrt{B_0 B_z}, \\ n_s &= n_s m_i / (\varepsilon_0 B_0 B_z), \\ T_s &= T_s / (m_e c^2). \end{aligned} \quad (11)$$

We also define

$$\begin{aligned} j_s &= n_s \frac{v_s}{c} = -\frac{\mu_0 c}{\Omega_i B_z} j_s, \\ N_s &= n_s - n_{0s}, \\ F &= \frac{F_y}{en_i c \sqrt{B_0 B_z}}. \end{aligned} \quad (12)$$

We can then collect equations (2), (3), (4), and (8) in the form

$$\begin{aligned} \partial_t E_x &= -A^2 \partial_z B - (1 - A^2) F, \\ \partial_t E_z &= \frac{B_z}{B_0} (\partial_x B + j_M + j_I), \\ \partial_t B &= \partial_x E_z - \partial_z E_x, \\ \partial_t N_s &= -\partial_z j_s, \\ \partial_t j_s &= -\frac{m_i}{m_e} n_s E_z - \partial_z \left(N_s T_s + \frac{j_s^2}{n_s} \right) + N_s \frac{\partial_z n_{0s} T_s}{n_{0s}}. \end{aligned} \quad (13)$$

Here it may be useful to point out some physical processes and how they are described by equations (13). Eliminating B between the first and third equations of (13), we obtain

$$(\partial_t^2 - A^2 \partial_z^2) E_x + A^2 \partial_z \partial_x E_z = -(1 - A^2) \partial_t F, \quad (14)$$

which describes driven shear Alfvén waves. While F is growing during the early stages of our simulations, the dynamics will mainly be determined by (14), and when F becomes stationary, the Alfvén waves are decoupled from the force. If we assume a stationary state and eliminate B from the first two equations of (13) we find, with $j = j_M + j_I$,

$$\partial_z j = -(1 - A^2) \partial_x F. \quad (15)$$

Equation (15) describes an MHD generator. The shear in the mechanical force generates a field-aligned current. During the later stages of our simulations the current from (15) determines E_z via the momentum equations, and in the electrostatic limit, $\partial_z E_x = \partial_x E_z$ gives the perpendicular electric field.

The equations of (13) describe the propagation of shear Alfvén waves and the formation of electrostatic structures but also high-frequency waves such as electrostatic oscillations at the plasma frequency ω_p and electromagnetic radiation. These high-frequency modes will cause severe problems if we try to integrate the equations with a time step $\Delta t \gg \omega_p^{-1}$ and a perpendicular resolution $\Delta x \ll c \Delta t$, suitable for the slow, short-scale processes we are interested in. In related numerical models [e.g., *Goertz and Boswell*, 1979; *Lysak and Dum*, 1983; *Streltsov and Lotko*, 1999] this problem is handled by neglecting the parallel displacement current, $\varepsilon_0 \partial_t E_z$. If we in Ampere's law replace ε_0 by a dielectric tensor according to

$$\varepsilon_0 \begin{pmatrix} 1 & 0 & 0 \\ 0 & 1 & 0 \\ 0 & 0 & \varepsilon_z \end{pmatrix}, \quad (16)$$

where $\varepsilon_z = 1$ in a normal vacuum, the second equation in (13) becomes

$$\partial_t E_z = \frac{B_z}{\varepsilon_z B_0} (\partial_x B + j_M + j_I). \quad (17)$$

We see that neglecting the parallel displacement current corresponds to $\varepsilon_z \rightarrow 0$, or $\omega_{pz}^2 = \omega_p^2 / \varepsilon_z \rightarrow \infty$, which means that the current is assumed to adjust instantaneously to any changes in the magnetic field. When B_y varies on a perpendicular scale comparable to the electron inertial length $\lambda_i = c / \omega_p$, it becomes difficult to justify this approximation, since the time it takes to build up a strong current increases as the plasma density decreases.

Rather than neglecting the displacement current by letting $\omega_{pz} \rightarrow \infty$, we will here use a less drastic method. We will instead increase the vacuum polarizability ε_z and slow down the plasma oscillations to a frequency that permits the numerical integration to converge. We choose

$$\varepsilon_z = \max(\omega_p \Delta t, 1), \quad (18)$$

which allows us to resolve the normal plasma oscillations whenever $\omega_p \Delta t \lesssim 1$. When $\omega_p \Delta t > 1$, the frequency of the plasma oscillations is brought down to $\omega_{pz} \approx \Delta t^{-1}$, which can be handled by the numerical integration. As long as the timescales we are interested in are well resolved by our Δt , it should not matter whether the plasma frequency is ω_p or ω_{pz} since they are both very high compared to the frequencies of interest. Notice that the phase velocity of electromagnetic radiation in the ordinary mode is also reduced to $c_z = c / \sqrt{\varepsilon_z}$. This means that the perpendicular resolution is improved to $\Delta x \sim c_z \Delta t$, which is comparable to $\lambda_i = c / \omega_p = c_z / \omega_{pz}$. When an anisotropic vacuum is introduced with a dielectric constant defined as in (18), the set of equations (13) becomes very convenient for numerical integration. These equations are integrated by a fully implicit method. Further details about the numerical method are given in Appendix A.

2.2. Model Parameters and Boundary Conditions

Although the model contains several parameters that can be varied, we will use a single set of values for most of them in this first study. We simulate a region $-L_x < x < L_x$, where L_x is 500 km at the generator and 21.1 km in the ionosphere. The length of the simulation box is $L_z = 3 \times 10^4$ km, which is a few R_E less than the distance from the ionosphere to the equatorial plane along a typical auroral field line. Since we neglect kinetic effects, the dispersion of short perpendicular wavelength Alfvén waves at higher altitudes is not described in detail by our model [*Lysak and Lotko*, 1996]. This shortening will lead to slightly shorter particle and Alfvén wave transit times but has little effect on the overall physics. Shortening the system saves

computer time and makes the interesting parts of the resulting plots easier to read by minimizing the featureless section between the generator and the acceleration region.

The magnetic field strength along an auroral field line is rather well approximated by

$$B_z(z) = B_0 \exp \left\{ z^2 / [L_z^2 \ln(B_I/B_0)] \right\}, \quad (19)$$

where we will use the ionospheric field strength $B_I = 56 \mu\text{T}$ and $B_0 = 0.1 \mu\text{T}$. This magnetic field model, which is justified by noticing that a logarithmic plot of the dipolar field strength on auroral latitudes versus z resembles a parabola, implies that the distance between field lines increases by $\sqrt{B_I/B_0} \approx 23.6$ from the ionosphere to the generator.

The force F_y applied in the generator region has the form

$$F_y(x, z, t) = F_0 \frac{\exp(-z^2/L_{zG}^2)}{\cosh^2(x/L_{xG})} \frac{t^2}{t^2 + t_G^2}, \quad (20)$$

where $L_{zG} = 6 \times 10^3$ km and $L_{xG} = 50$ km. The maximum force F_0 will be 1.5×10^{-17} N m $^{-3}$, and in this study we will use the generator timescale $t_G = 10$ s.

The density of the magnetospheric electrons is initially set to

$$n_{0M}(z) = \left[1 + \frac{B_z}{B_I} \left(\frac{B_z}{B_0} \right)^{\frac{1}{4}} \right] 3 \times 10^5 \text{ m}^{-3}, \quad (21)$$

which is roughly constant down to below the acceleration region and then increases to $1.76 \times 10^6 \text{ m}^{-3}$ near the ionosphere. The temperature of the magnetospheric electron component is kept constant at $T_M = 1$ keV.

In terms of the height $h = L_z - z$ above the lower boundary, the density of ionospheric electrons is given by

$$n_{0I}(h) = \left\{ \exp \left[\frac{-h}{H(1 + h/5R_E)} \right] + 10^{-8} \right\} 10^9 \text{ m}^{-3}. \quad (22)$$

Here, $H = 400$ km is the scale height in the ionosphere, but the effective scale height increases slightly with altitude. The temperature of the ionospheric electrons is set to $T_I(h) = 0.1 \text{ eV} \sqrt{n_{0I}(0)/n_{0I}(h)}$, with an upper limit of 100 eV.

The lower ionosphere is represented by a height-integrated ionospheric Pedersen conductivity, which we take to be $\Sigma_P = 10 \Omega^{-1}$. Integrating Ohm's law and Ampere's law across the ionospheric height leads to the condition

$$B_y(L_z) = \mu_0 \Sigma_P E_x(L_z) \quad (23)$$

at the ionospheric boundary. For precipitating electrons ($j_s > 0$) we demand $\partial_z j_s = 0$ at $z = L_z$, which also implies $\partial_t N_s(L_z) = 0$. However, we do not allow elec-

trons to flow up from the boundary into the simulation box, since that would add momentum to the system in an uncontrolled way. Hence we set $j_s(L_z) = 0$ and $\partial_z N_s(L_z) = 0$ when $j_s < 0$ inside the boundary. No ionospheric boundary condition is needed for E_z .

Since the force F_y is symmetric around the equator, we assume that E_x and the densities n_s also are symmetric. Their boundary conditions at $z = 0$ are thus $\partial_z E_x = \partial_z N_s = 0$. The remaining fields are then anti-symmetric, so we set $E_z = B = j_s = 0$ at the equatorial plane. The fields are all set to zero on the boundaries at $x = \pm L_x$.

3. Results

To illustrate the response of the auroral plasma to a force applied in the equatorial region, we will show plots of selected fields at various times. The force F_y is given by (20) with $F_0 = 1.5 \times 10^{-17}$ N m $^{-3}$ and $t_G = 10$ s. When this force starts to pull the equatorial plasma in the y direction, the divergence of the perpendicular current causes a charge separation that sets up an electric field (E_x, E_z). The growth of E_x and the current j_z driven by E_z are coupled via the magnetic perturbation B_y .

Figures 2a–2d show perspective plots of these fields 1 s after the force has been turned on. The fields in physical units are plotted as functions of the field-aligned coordinates (x, h) , and the distances given on the x axis refer to the ionospheric boundary ($h = 0$). We show only $|x| < 10$ km, although the simulation covers ± 21 km. The perpendicular electric field is still confined to the equatorial region and is roughly proportional to F_y . There are peaks in B_y and E_z at the altitude where $\partial_z E_x$ is large. A small E_z is seen at altitudes around 7200 km, where the Alfvén velocity reaches its maximum value of $\sim 1.1 \times 10^8$ m s $^{-1}$. A significant current density, which causes a noticeable magnetic perturbation, is already starting to build up near the ionospheric boundary. Considering that the time it takes an Alfvén wave to propagate through the system,

$$t_{L_z} = \int_0^{L_z} \frac{dz}{cA}, \quad (24)$$

is about 3.6 s, it may be surprising that information about the generator has already reached the ionosphere. However, signals can also propagate with the sound speed 1.87×10^7 m s $^{-1}$ of the magnetospheric electrons. By combining the two modes of propagation, a precursor can reach the ionosphere in ~ 1 s. Other simulations ran with $t_G = 0.1$ s show that the main pulse reaches the ionosphere at $t \sim t_{L_z}$ as expected.

In Figures 3a–3d we show the fields after 10 s, when the force has grown to half its final value. The fields are now much more concentrated to the lower part of the flux tube. In particular, the magnetic perturbation B_y and the field-aligned current density $j_z = j_M + j_I$

reach their maxima at the ionospheric boundary where the background magnetic field is strongest. The perpendicular electric field has a more complicated structure, which results from the combination of the growing pulse coming down from the generator and the reflected pulse from the ionosphere. The field-aligned electric field is concentrated to altitudes around 7000 km, which is roughly where the normalized density $n = n_M + n_I$ has its minimum.

After 60 s the generator force has reached 97.3 % of its final strength, and the fields shown in Figures 4a–4d have become rather stationary. The magnetic disturbance B_y is roughly proportional to $\sqrt{B_z}$, and j_z is proportional to B_z . This is what we expect, since (13) indicates that $\partial_z \mathbf{B} = \partial_z \mathbf{j}_s = 0$ below the generator in a stationary state. The strongest E_x are found near the acceleration region, between the peaks in E_z . Notice the ridges with $E_x > 0$ at high altitudes, which indicate regions where the plasma flow is in a direction opposite to the applied force. In the upward current region there is a strong negative E_z accelerating the precipitating magnetospheric electrons, but in the downward current region the sign of E_z changes with altitude. At lower altitudes a positive E_z accelerates electrons upward from the ionosphere. Above the acceleration region a nega-

tive E_z is required to slow the electrons down as the flux tube widens and n increases.

A persistent feature in our simulations is that the downward currents are narrower and more intense than the upward currents, as seen in Figure 4d. Even if we at present cannot explain the reasons for this difference, it is interesting to notice that a similar effect was observed by the Fast Auroral SnapshoT (FAST) satellite [Elpic *et al.*, 1998].

Although our simulations are based on the electromagnetic fields, it is instructive to consider also the potentials. In our case there are two straightforward ways to calculate the scalar potential. Setting the potential to zero in the ionosphere at $x = -L_x$, we can integrate the electric field to obtain

$$\phi_z(x, z, t) = - \int_{-L_x}^x E_x(x', L_z) dx' - \int_{L_z}^z E_z(x, z') dz'. \quad (25)$$

The scalar potential ϕ_z corresponds to a particular choice of gauge, in which the vector potential $\mathbf{A} = A_x \hat{x}$ and the fields are given by $E_x = -\partial_x \phi_z - \partial_t A_x$, $E_z = -\partial_z \phi_z$, and $B_y = \partial_z A_x$. Alternatively, we can also compute the potential by doing the integrals in the reverse order as

Figure 2. Perspective plots 1 s after the generator was started of (a) the magnetic field perturbation, (b) the parallel electric field, (c) the perpendicular electric field, and (d) the parallel current density.

Figure 3. Perspective plots 10 s after the generator was started of (a) the magnetic field perturbation, (b) the parallel electric field, (c) the perpendicular electric field, and (d) the parallel current density.

$$\phi_x(x, z, t) = - \int_{L_z}^z E_z(-L_x, z') dz' - \int_{-L_x}^x E_x(x', z) dx', \quad (26)$$

and in this gauge the fields are $E_x = -\partial_x \phi_x$, $E_z = -\partial_z \phi_x - \partial_t A_z$, and $B_y = -\partial_x A_z$. In the stationary limit these two scalar potentials will become equal, but they include different parts of the time variations.

Figure 5a shows the shape of ϕ_z at $t = 90$ s. The general shape of ϕ_z becomes similar to this a few seconds after the start, and it retains this shape while the magnitude grows. At $t = 90$ s the shape of ϕ_x shown in Figure 5b is rather similar to that of ϕ_z , indicating that the field is mainly electrostatic. However, ϕ_x also contains a significant oscillating component. Figure 6 shows $\phi_x(x, 0, t)$ in the equatorial plane at times $t = 71, 76, 81$, and 86 s. We interpret this oscillation as a field line resonance, a standing shear Alfvén wave with period $t_R \approx 20$ s that is excited by the growing force and then decouples from the generator when F_y becomes stationary. The period t_R remains constant when t_G is varied, which confirms the resonant character of the oscillation. The time-dependent electric field associated with this Alfvén resonance will be mainly in the x direction, and depending on the gauge chosen, it can be seen in either A_x or ϕ_x , but hardly at all in ϕ_z .

The collisionless electrons in the auroral magnetosphere are in this study modeled by compressible fluids, and the acceleration of a fluid is in some respects different from the acceleration of noninteracting particles. Below the acceleration region, where the parallel electric field is short-circuited by the ionospheric plasma, the magnetospheric electron fluid must be accelerated by pressure gradients. Figure 7 shows the density perturbation profile at $t = 90$ s along a field line that reaches the ionosphere at $x = -1.3$ km, which is in the middle of the upward current region. As shown in Figure 7, the density of magnetospheric electrons piles up in the lower part of the acceleration region. The electrons are then accelerated to their final energy by the gradient below the density maximum. The final energy of the precipitating electrons depends on the density of the magnetospheric electrons at the ionospheric boundary. Figure 8 shows the kinetic energy of the magnetospheric electrons along the field line at $x = -1.3$ km at $t = 90$ s. With the model we use for n_{0M} the kinetic energy reached at the ionospheric boundary is very similar to the energy $e\Delta\phi$ gained by falling through the potential drop $\Delta\phi$ along the field line (compare Figure 5). We emphasize that although the compressibility of the magnetospheric electron fluid complicates the picture of the acceleration process, the potential drop

Figure 4. Perspective plots 60 s after the generator was started of (a) the magnetic field perturbation, (b) the parallel electric field, (c) the perpendicular electric field, and (d) the parallel current density.

in our simulations is supported by electron inertia and the low plasma density above the auroral ionosphere as discussed by Rönmark [1999].

The density of ionospheric electrons is strongly depleted at altitudes around 9000 km in the upward current region. To prevent numerical problems with neg-

ative densities (see Appendix A), we must limit this density reduction to $\sim 80\%$. A density perturbation profile for the ionospheric electrons along the field line at $x = -1.3$ km at $t = 90$ s is shown in Figure 9.

In the downward current region the density of ionospheric electrons grows and builds up to a sharp peak

Figure 5. Perspective plots 90 s after the generator was started of (a) the scalar potential ϕ_z and (b) ϕ_x .

Figure 6. Cross section of the scalar potential ϕ_x in the equatorial plane at four times. The oscillation indicates that a field line resonance with a period $t_R \sim 20$ s is excited.

Figure 8. The kinetic energy of magnetospheric electrons along a field line at $x = -1.3$ km at $t = 90$ s.

near the reversal of E_z at $h = 8500$ km. This peak continues to move slowly upward throughout the simulation, while N_I grows to almost $100 n_{0I}$. The separation of the electrons into two fluids is not so easily justified in this region, and we are not convinced that the physics is described in detail by our model.

The total potential drop along a field line, given by $\Delta\phi(x, t) = \phi_z(x, L_z, t) - \phi_z(x, 0, t)$, is in Figure 10 plotted versus the current density $j_M(x, L_z, t)$ carried

into the ionosphere by the magnetospheric electrons. This is the current-voltage relation a spacecraft passing through the upward current region would observe at $t = 90$ s. The crosses show values from the simulation, and a solid line shows the theoretical current-voltage relation derived by Rönmark [1999]

$$\Delta\phi(x, t) = \frac{m_e}{2e^3 n^2} j_M^2. \quad (27)$$

Figure 7. The relative density perturbation N_M/n_{0M} in the magnetospheric electrons along a field line at $x = -1.3$ km at $t = 90$ s.

Figure 9. The relative density perturbation N_I/n_{0I} in the ionospheric electrons along a field line at $x = -1.3$ km at $t = 90$ s.

Figure 10. The current-voltage relation in the upward current region. The crosses show the simulation results, and the solid line is the theoretical result predicted by equation (27).

Using the density $n = 1.76 \times 10^6 \text{ m}^{-3}$ for the magnetospheric electrons, we find $m_e/(2e^3n^2) = 3.576 \times 10^{13} \text{ V A}^{-2}$, while a least squares fit to the crosses gives $3.593 \cdot 10^{13} \text{ V A}^{-2}$. The agreement between the electrostatic fluid model and our numerical simulations is almost perfect, and our simulations strongly support that the self-consistent potential drop on auroral field lines is proportional to the square of the current density.

4. Discussion

The model developed in this study includes a generator that converts work done by equatorial, mechanical forces into electromagnetic power. The propagation of electromagnetic energy down to the acceleration region is described both during the dynamic stage dominated by Alfvén waves and during the electrostatic stage. In the acceleration region, electromagnetic power is converted back to mechanical energy in the form of an intense flux of precipitating energetic electrons.

There are several mechanisms that contribute to the appearance of parallel electric fields in shear Alfvén waves. At altitudes above $\sim 4 R_E$, parallel electric fields can be caused by finite Larmor radius effects [Hasegawa, 1976]. As discussed in section 2, these effects are not included in our model, but this omission is not likely to be important since most of the electron acceleration is expected to occur at lower altitudes. At these lower altitudes an E_z is generated by electron inertia, that is, by terms proportional to the electron mass in the momentum equation

$$d_t m_e v_z = \partial_t m_e v_z + m_e v_z \partial_z v_z \approx -e E_z. \quad (28)$$

In some studies [e.g., Goertz and Boswell, 1979; Streltsov and Lotko, 1999; Streltsov, 1999] only the partial time derivative $\partial_t m_e v_z$ is included, but the convective term is far more important in the auroral acceleration region for the slow, large-amplitude processes we consider. The convective term is included by Lysak and Dum [1983], but its effects are not seen since n/B is constant in their model. As discussed by Rönnmark [1999], the term $m_e v_z \partial_z v_z$ alone can account for stationary, electrostatic potential drops of several kilovolts when the variation of n/B along auroral field lines is taken into account. The importance of the convective term for electron acceleration by field line resonances has also recently been stressed by A. N. Wright et al. (Parallel electric fields in the auroral acceleration region of field line resonances, submitted to *Journal of Geophysical Research*, 2000).

When describing the generator, we have assumed that the force F_y is independent of the convection velocity $u_y = -E_x/B_z$ reached by the equatorial plasma. This corresponds to a pure current generator, and this idealized case was chosen for its simplicity. Mixtures between current and voltage generators can be modeled by allowing the force F_y to depend on u_y , but it is not obvious how this dependence should be chosen. A self-consistent treatment of the generator probably requires a three-dimensional simulation.

In the simulations presented here the auroral current sheets are $\sim 1 \text{ km}$ thick and separated by $\sim 2 \text{ km}$ at ionospheric altitude. These scales, which are determined by our choice of F_y , are typical for discrete arc systems, which may split up into several finer scale arcs when observed optically from the ground [Borovsky, 1993]. Our simulations show that the ionospheric resistance, which is essential for the large-scale magnetosphere-ionosphere coupling, is almost negligible for such small structures. This is also easily confirmed by an order of magnitude estimate, which shows that even if the ionospheric conductivity is as low as $1 \Omega^{-1}$ the potential drop in the ionosphere will be just a few volts for an auroral current density of $3 \mu\text{A m}^{-2}$ that yields a field-aligned potential drop $\sim 350 \text{ V}$. This makes it very unlikely that enhancements of the ionospheric conductivity can have significant effects on the development of small-scale auroral structures.

The fluid model we use cannot fully describe the behavior of a collisionless plasma. As mentioned in section 3, the magnetospheric electrons must be accelerated by pressure gradients below the acceleration region. This limits the velocity they can reach to the sound velocity, which in our isothermal model equals their thermal velocity. The present version of the code behaves well for potential drops less than $\sim 400 \text{ V}$, but for higher potentials it becomes difficult to control the density fluctuations.

The results presented in Figure 10 strongly support that the potential drop on auroral field lines is proportional to the square of the current. This forces

us to reconsider the observations that have been interpreted as supporting a linear current-voltage relationship [Lyons *et al.*, 1979; Weimer *et al.*, 1987; Lu *et al.*, 1991; Haerendel *et al.*, 1994]:

$$j_z \approx -K \Delta\phi. \quad (29)$$

These are event studies, and they contain no statistical evidence indicating whether the data are best fitted by a linear relation or some other function. In these studies the Lyons-Evans-Lundin constant K [Fridman and Lemaire, 1980] is a free parameter, which varies by almost 2 orders of magnitude from event to event (and sometimes is adjusted during an event). Considering that the variation of the observed potential drop in each event is rather limited, it is probably possible to find a reasonable linear fit even if $\Delta\phi$ really is proportional to j_z^2 .

In a recent study of 22 events [Sakanoi *et al.*, 1995] the latitudinal distribution of the upward current density was found to have an inverted-U shape, rather than the inverted-V shape of the potential. In many cases, $\Delta\phi/j_z$ was ~ 5 –10 times larger at the center of the inverted-V potential than at the edges. Assuming that the plasma density is independent of latitude during each event, this result is expected if the potential drop is proportional to the square of the current. In our opinion, this study indicates that the observational support for a linear current-voltage relationship is rather weak.

There are also reasons to question the theoretical arguments in favor of a linear current-voltage relation. The kinetic models that predict a linear relation are not self-consistent but based on the adiabatic orbits of test particles. Initially, the particle distributions at the equatorial and ionospheric boundaries are specified in a state which we may assume to be charge neutral and without current. The adiabatic particle orbits will determine the distribution function, and hence the density, along the flux tube. An essentially arbitrary electrostatic potential is then prescribed, and new particle orbits are calculated by taking this potential into account. The particle distributions are then integrated to obtain the current, and the density can also be calculated. If the test particles in the model were real electrons, the densities calculated from the new orbits would imply large deviations from charge neutrality. Jauhunen [1999] used this method to calculate the density in the presence of an upward electric field, and he found that for $e\Delta\phi \sim 2T_M$ the density increase at low altitudes was comparable to the total equatorial density. A corresponding increase in the number of real electrons would cause a strong downward electric field, which is not included in the calculation of the particle orbits. As long as the particle distribution function calculated from the specified boundary conditions produces a charge density that is inconsistent with the potential, we have no reason to expect it to predict the correct current.

Even if kinetic test particle models cannot be used to obtain the current-voltage relation, they provide valuable information about the behavior of suprathermal particles in the auroral acceleration region. As long as the high-energy particles constitute a small fraction of the density, they will behave approximately as predicted by these models. In particular, the accelerated electrons below the parallel electric field will behave as test particles. Therefore the theoretical studies of velocity space boundaries between different particle populations, and the observations that confirm these theories, are not much affected by the lack of self-consistency. Some examples of such studies are discussed in the review by Lundin and Eliasson [1991].

In the literature, auroral electron acceleration by an electrostatic field has usually been treated as entirely separated from acceleration by shear Alfvén waves. Speaking about Alfvén wave models, Lysak and Dum [1983, p. 365] stated that “One problem with such models is that they do not possess a steady state parallel electric field to accelerate particles.” and it seems that this opinion implicitly has been shared by most workers in this field. However, our simulations demonstrate that shear Alfvén waves build up an electric field that becomes electrostatic when the driving force is stationary. Even in the absence of collisions, there is a smooth transition from driven Alfvén waves to an electrostatic field maintained by a magnetospheric generator.

Appendix A: Numerical Method

The set of equations in (13) is solved by a fully implicit method, based on factorization of the two-dimensional spatial differential operators. Algorithms of this type are discussed by, for example, Degrez [1992]. If we for notational convenience collect all the fields in a vector $\mathbf{U} = (E_x, E_z, B_y, N_s, j_s)^T$, we can after time discretization sum up (13) in the form

$$\mathbf{U}(t + \Delta t) = \mathbf{U}(t) + \Delta t \{ \mathbf{A}(\mathbf{U}(t + \Delta t/2)) + \mathbf{G} \}, \quad (30)$$

where $\mathbf{A}(\mathbf{U})$ represents the $\mathbf{U}$ dependence of the right hand side of (13) and $\mathbf{G}$ represents some inhomogeneous terms involving the external force $\mathbf{F}$ and the densities n_{0s} . We now linearize $\mathbf{A}$ in $\mathbf{U}(t + \Delta t)$ by a Taylor expansion:

$$\begin{aligned} & \mathbf{A}(\mathbf{U}(t + \Delta t/2)) \\ & \approx \mathbf{A}(\mathbf{U}(t)) + \frac{1}{2} [\mathbf{U}(t + \Delta t) - \mathbf{U}(t)] \cdot \partial_{\mathbf{U}} \mathbf{A}(\mathbf{U}(t)) \\ & = \frac{1}{2} [\mathbf{U}(t + \Delta t) + \mathbf{U}(t)] \cdot \partial_{\mathbf{U}} \mathbf{A}(\mathbf{U}(t)), \end{aligned} \quad (31)$$

where the second step follows from the homogeneous properties of $\mathbf{A}$. This results in a linear set of equations:

$$\left[\mathbf{I} - \frac{\Delta t}{2} \partial_{\mathbf{U}} \mathbf{A} \right] \cdot \mathbf{U}(t + \Delta t) = \left[\mathbf{I} + \frac{\Delta t}{2} \partial_{\mathbf{U}} \mathbf{A} \right] \cdot \mathbf{U}(t) + \Delta t \mathbf{G}. \quad (32)$$

The operator $\partial_U \mathbf{A}$, which contains both ∂_x and ∂_z , is now split into two parts as $\partial_U \mathbf{A} = \mathbf{X} + \mathbf{Z}$, where $\mathbf{X}$ contains only ∂_x and $\mathbf{Z}$ contains only ∂_z . This decomposition is not unique, and it should be chosen in a way that makes the product $\mathbf{X} \cdot \mathbf{Z}$ as small and simple as possible. Neglecting a small term $\Delta t^2/4 \mathbf{X} \cdot \mathbf{Z} \cdot [\mathbf{U}(t + \Delta t) - \mathbf{U}(t)]$, we then find

$$\begin{aligned} \left[\mathbf{I} - \frac{\Delta t}{2} \mathbf{X} \right] \cdot \left[\mathbf{I} - \frac{\Delta t}{2} \mathbf{Z} \right] \cdot \mathbf{U}(t + \Delta t) = \\ \left[\mathbf{I} + \frac{\Delta t}{2} \mathbf{X} \right] \cdot \left[\mathbf{I} + \frac{\Delta t}{2} \mathbf{Z} \right] \cdot \mathbf{U}(t) + \Delta t \mathbf{G}. \end{aligned} \quad (33)$$

Introducing $\mathbf{U}^* = [\mathbf{I} - \Delta t/2 \mathbf{Z}] \cdot \mathbf{U}(t + \Delta t)$ as a new variable, we can solve (33) in two steps. First we solve

$$\begin{aligned} \left[\mathbf{I} - \frac{\Delta t}{2} \mathbf{X} \right] \cdot \mathbf{U}^* = \\ \left[\mathbf{I} + \frac{\Delta t}{2} \mathbf{X} \right] \cdot \left[\mathbf{I} + \frac{\Delta t}{2} \mathbf{Z} \right] \cdot \mathbf{U}(t) + \Delta t \mathbf{G}. \end{aligned} \quad (34)$$

for $\mathbf{U}^*$. Then we use

$$\left[\mathbf{I} - \frac{\Delta t}{2} \mathbf{Z} \right] \cdot \mathbf{U}(t + \Delta t) = \mathbf{U}^* \quad (35)$$

to solve for the fields at $t + \Delta t$. When the operators $\mathbf{X}$ and $\mathbf{Z}$ are expressed as centered finite differences, each of these two steps consists of solving a block-tridiagonal set of equations. The number of operations needed for this scales linearly with the mesh size ($N_x \times N_z$) and hardly requires any extra memory, which makes this algorithm very efficient. The number of operations needed for a direct integration of (32) would typically be proportional to $(N_x \times N_z)^2$.

The simulations presented here were all performed on a rectangular, slightly nonuniform grid of size $(N_x \times N_z) = (51 \times 120)$, and we used a time step $\Delta t = 0.02$ s.

The algorithm described above is numerically stable, but the grid tends to become decoupled in the ionosphere where E_z is small. To cure this, even and odd grid points are coupled by averaging the fields E_x , n_s , and j_s as

$$\mathbf{U}(z) = 0.9\mathbf{U}(z) + 0.05[\mathbf{U}(z + \Delta z) + \mathbf{U}(z - \Delta z)]. \quad (36)$$

Some extra precautions are needed to ensure that the code behaves well in simulations that should converge to a quasi-stationary state with potential drops of hundreds of volts. To prevent the densities from going to zero near boundaries where the current is forced to vanish, we add a source term of the form $-S(z)N_s$ to the continuity equations. The function $S(z)$ is exponentially small away from the boundaries, except that it in the upward current region includes a term $10^{-3}(n_{01}/n_1)^3$ which becomes significant for very small n_1 . If the ionospheric electrons are allowed to become supersonic, steep density gradients will form in the lower part of the acceleration region, where their density already is reduced by typically 80 %. To prevent this from making

the density negative, we transfer some momentum from ionospheric to magnetospheric electrons by subtracting a force $\propto 0.05 m_e v_I^2 / T_I$ in the momentum equation for the ionospheric electrons, and we conserve total current by adding a corresponding term to the magnetospheric component.

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